

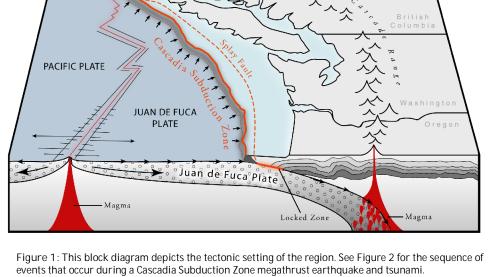
## Introduction

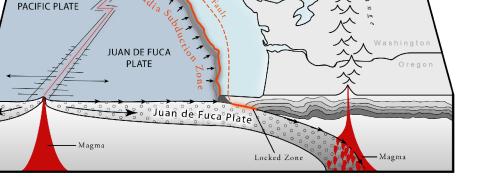
CSZ Model Specifications. The sizes of the earthquake and its resultant tsunami are primarily driven The Oregon Department of Geology and Mineral Industries (DOGAMI) has been identifying and mapping the tsunami inundation hazard along the Oregon coast since 1994. In Oregon, DOGAMI by the amount and geometry of the slip that takes place when the North American Plate snaps manages the National Tsunami Hazard Mitigation Program, which has been administered by the westward over the Juan de Fuca Plate during a CSZ event. DOGAMI has modeled a wide range of to establish the area of expected tsunami inundation based on scientific evidence and tsunami National Oceanic and Atmospheric Administration (NOAA) since 1995. DOGAMI's work is designed earthquake and tsunami sizes that take into account different fault geometries that could amplify to help cities, counties, and other sites in coastal areas reduce the potential for disastrous tsunamithe amount of seawater displacement and increase tsunami inundation. Seismic geophysical related consequences by understanding and mitigating this geologic hazard. Using federal funding profiles show that there may be a steep splay fault running nearly parallel to the CSZ but closer to the Oregon coastline (Figure 1). The effect of this splay fault moving during a full-rupture CSZ event awarded by NOAA, DOGAMI has developed a new generation of tsunami inundation maps to help residents and visitors along the entire Oregon coast prepare for the next Cascadia Subduction Zone would be an increase in the amount of vertical displacement of the Pacific Ocean, resulting in an (CSZ) earthquake and tsunami. increase of the tsunami inundation onshore in Oregon. DOGAMI has also incorporated physical evidence that suggests that portions of the coast may drop 4 to 10 feet during the earthquake; this The CSZ is the tectonic plate boundary between the North American Plate and the Juan de Fuca effect is known as subsidence. Detailed information on fault geometries, subsidence, computer

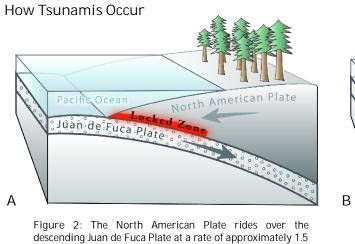
Plate (Figure 1). These plates are converging at a rate of about 1.5 inches per year, but the movement is not smooth and continuous. Rather, the plates lock in place, and unreleased energy builds over time. At intervals, this accumulated energy is violently released in the form of a megathrust earthquake rupture, where the North American Plate suddenly slips westward over the Juan de Fuca Plate. This rupture causes a vertical displacement of water that creates a tsunami (Figure 2). Similar rupture processes and tsunamis have occurred elsewhere on the planet where subduction zones exist: for example, offshore Chile in 1960 and 2010, offshore Alaska in 1964, near Sumatra in 2004, and offshore Japan in March 2011.

CSZ Frequency. Comprehensive research of the offshore geologic record indicates that at least 19 major ruptures of the full length of the CSZ have occurred off the Oregon coast over the past 10,000 years (Figure 3). All 19 of these full-rupture CSZ events were likely magnitude 8.9 to 9.2 earthquakes (Witter and others, 2011). The most recent CSZ event happened 312 years ago on January 26, 1700. Sand deposits carried onshore and left by the 1700 event have been found 1.2 miles inland; older tsunami sand deposits have also been discovered in estuaries 6 miles inland. As shown in Figure 3, the range in time between these 19 events varies from 110 to 1,150 years, with a median time interval of 490 years. In 2008 the United States Geological Survey (USGS) released the results of a study announcing that the probability of a magnitude 8-9 CSZ earthquake occurring over the next 30 years is 10% and that such earthquakes occur about every 500 years (WGCEP, 2008).

## Cascadia Subduction Zone Setting JUAN DE FUCA







cumulative number of buildings inundated within the map area.

models, and the methodology used to create the tsunami scenarios presented on this map can be

found in DOGAMI Special Papers 41 (Priest and others, 2009) and 43 (Witter and others, 2011).

This tsunami inundation map displays the output of computer models representing five selected

tsunami scenarios, all of which include the earthquake-produced subsidence and the tsunami-

amplifying effects of the splay fault. Each scenario assumes that a tsunami occurs at Mean Higher

High Water (MHHW) tide; MHHW is defined as the average height of the higher high tides observed

material and to enhance the educational aspects of hazard mitigation and response, the five

Extra Large (S, M, L, XL, XXL). The map legend depicts the respective amounts of slip, the frequency

of occurrence, and the earthquake magnitude for these five scenarios. Figure 4 shows the

The computer simulation model output is provided to DOGAMI as millions of points with values

that indicate whether the location of each point is wet or dry. These points are converted to wet and

dry contour lines that form the extent of inundation. The transition area between the wet and dry

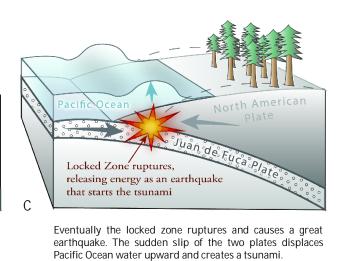
contour lines is termed the Wet/Dry Zone, which equates to the amount of error in the model when

determining the maximum inundation for the each scenario. Only the XXL Wet/Dry Zone is shown

scenarios are labeled as "T-shirt sizes" ranging from Small, Medium, Large, Extra Large, to Extra

over an 18-year period at the Port Orford tide gauge. To make it easier to understand this scientific

Because the two plates are stuck in place at the "locked zone," strain builds up over time and the North American



Research-indicated radiocarbon age of CSZ event (most recent in January 1700)

Figure 3: This chart depicts the timing, frequency, and magnitude of the last 19 great Cascadia

Subduction Zone events over the past 10,000 years. The most recent event occurred on January 26,

1700. The 1700 event is considered to be a "medium sized" event. The data used to create this chart

came from research that examined the many submarine landslides, known as "turbidites," that are

This map also shows the regulatory tsunami inundation line (Oregon Revised Statutes 455.446 and

455.447), commonly known as the Senate Bill 379 line. Senate Bill 379 (1995) instructed DOGAMI

modeling in order to prohibit the construction of new essential and special occupancy structures in

Time Series Graphs and Wave Elevation Profiles. In addition to the tsunami scenarios, the computer

model produces time series data for "gauge" locations in the area. These points are simulated gauge

stations that record the time, in seconds, of the tsunami wave arrival and the wave height observed

for each time interval. It is especially noteworthy that the greatest wave height and velocity observed

are not necessarily associated with the first tsunami wave to arrive onshore. Therefore evacuees

should not assume that the tsunami event is over until the proper authorities have sounded the all-

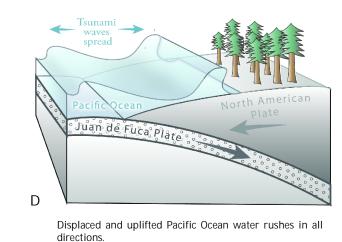
clear signal at the end of the evacuation. Figure 5 depicts time series data for the map plate area.

Figure 6 (profiles A-A' and B-B') depicts the overall wave height and inundation extent for all five

Occurrence and Relative Size of Cascadia Subduction Zone

this tsunami inundation zone, (Priest, 1995).

scenarios at select profiles on this map.

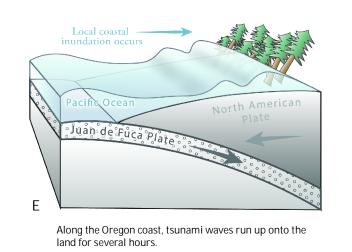


Buildings Inside Tsunami Zones \*

Percent of Buildings Inside Tsunami Zones

\*Building counts shown are based on polygon centroids and are cumulative.

scenario" for cities and unincorporated portions of the map.

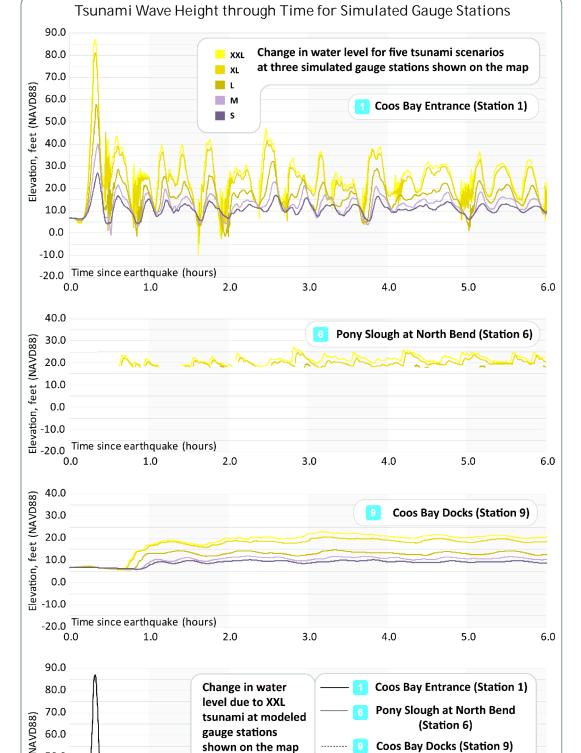


Coos Bay Area Buildings within Tsunami Inundation Zones

Extra Extra Large 21.9% 17.1% 14.4% 32.2%

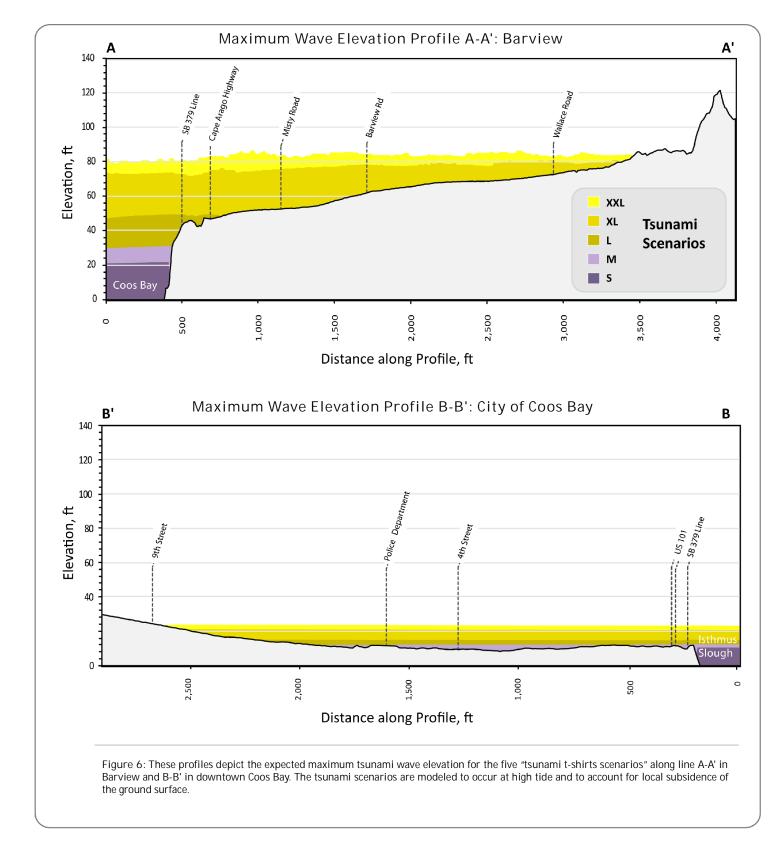
Entire Map Area City of Coos Bay City of North Bend Unincorporated Areas

Figure 4: The table and chart show the number of buildings inundated for each "tsunami T-shirt

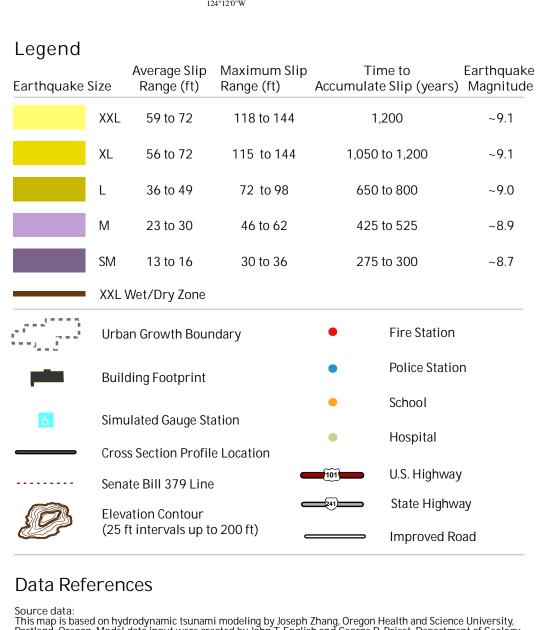


-20.0 Time since earthquake (hours)

1.0 2.0



◀ Tsunami Wave Height over Time for Simulated Gauge Stations Figure 5: Top three charts depict the tsunami waves as they arrive at the selected reference points (simulated gauge stations). It shows the change in wave heights for all five tsunami scenarios over an 6-hour period. The model predicts the first tsunami wave will arrive at the entrance to Coos Bay in approximately 20 minutes. The starting water elevation (0.0 hour) takes into account the local land subsidence or uplift caused by the earthquake. Wave heights vary through time, and the first wave will not necessarily be the largest as waves interfere and reflect off local topography and bathymetry. Bottom chart depicts the change in wave height for the XXL tsunami scenario only. Modeled wave heights, arrival times, and wave durations can help emergency response personnel plan for a tsunami. (Chart revised 07/15/2012.)

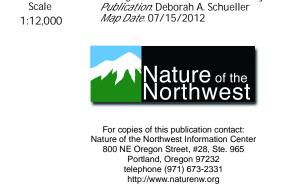


Source data:
This map is based on hydrodynamic tsunami modeling by Joseph Zhang, Oregon Health and Science University,
Portland, Oregon. Model data input were created by John T. English and George R. Priest, Department of Geology
and Mineral Industries (DOGAMI), Portland, Oregon. Hydrology data, contours, critical facilities, and building footprints were created by DOGAMI from 2009 to 201 Senate Bill 379 line data were redigitized by Rachel R. Lyles Smith and Sean G. Pickner, DOGAMI, in 2011 (GIS fi Urban growth boundaries (2010) were provided by the Oregon Department of Land Conservation and Development (DLCD). Transportation data (2008) were provided by Coos County.

Lidar data are from DOGAMI Lidar Data Quadrangles LDQ-2009-43124-C3-Charleston, LDQ-2009-43124-C2-CoosBay, LDQ-2009-43124-D2-NorthBend, and LDQ-2009-43124-D3-Empire. Coordinate System: Oregon Statewide Lambert Conformal Conic, Unit: International Feet, Datum: North American Datum 1983 HARN. Graticule shown with geographic coordinates (latitude/longitude).

References: 2007 Working Group on California Earthquake Probabilities (WGCEP), 2008, The Uniform California Earthquake Rupture Forecast, Version 2 (UCERF 2): U.S. Geological Survey Open-File Report 2007-1437 and California Geological Survey Special Report 203 [http://pubs.usgs.gov/of/2007/1437/]. Priest, G. R., 1995, Explanation of mapping methods and use of the tsunami hazard maps of the Oregon coast, Oregon Department of Geology and Minerals Industries Open-File Report O-95-67, 95 p. Priest, G.R., Goldfinger, C., Wang, K., Witter, R.C., Zhang, Y., and Baptista, A.M., 2009, Tsunami hazard assessment of the northern Oregon coast: a multi-deterministic approach tested at Cannon Beach, Clatsop County, Oregon: Oregon Department of Geology and Mineral Industries Special Paper 41, 87 p. Witter, R.C., Zhang, Y., Wang, K., Priest, G.R., Goldfinger, C., Stimely, L.L., English, J.T., and Ferro, P.A., 2011, Simulating tsunami inundation at Bandon, Coos County, Oregon, using hypothetical Cascadia and Alaska earthquake scenarios: Oregon Department of Geology and Mineral Industries Special Paper 43, 57 p.

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Coos-01 Lakeside West

Coos-02 Lakeside East

Coos-03 Saunders Lake Coos-04 Haynes Inlet

Coos-06 Coos River North

Coos-07 Coos River South

Coos-09 Barview - South Slough

Coos-05 Coos Bay - North Bend Coos-14 Coquille

Coos-08 Charleston - Cape Arago Coos-17 New River

0 0.075 0.15 0.3 Mile

Tsunami Inundation Map Index

Coos-10 Isthmus Slough

Coos-12 Bullards Beach

Coos-15 Coquille River

Map Data Creation/Development: *Tsunami Inundation Scenarios*: George R. Priest, Laura L. Stimely, Daniel E. Coe, Paul A. Ferro, Sean G. Pickner, Rachel R. Lyles Smith *Basemap Data*: Kaleena L.B. Hughes, Sean G. Pickner, Daniel E. Coe, Mathew A. Tilman

Cartography. Kaleena L.B. Hughes, Sean G. Pickner Text: Don W.T. Lewis, Rachel R. Lyles Smith Editing. Don W.T. Lewis, Rachel R. Lyles Smith