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Deep Landslide Susceptibility Map of the Bull Run Watershed

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Surficial and Bedrock Engineering Geology, Landslide Inventory and Susceptibility, and Surface Hydrography of the Bull Run Watershed, Clackamas and Multnomah Counties, Oregon

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PLATE 3

NOTICE

This product is for informational purposes and may not have been prepared for or be suitable for legal, engineering, or surveying purposes. Users of this information should review or consult the primary data and information sources to ascertain the usability of the information. This publication cannot substitute for site-specific investigations by qualified practitioners. Site-specific data may give results that differ from the results shown in the publication. See the accompanying text report for more details on the limitations of the methods and data used to prepare this publication.

ABOUT THIS PUBLICATION

The Bull Run Watershed (BRW) is the primary drinking water supply for the City of Portland and several suburbs and is cooperatively managed by the Portland Water Bureau and the U.S. Forest Service. The watershed is located 25 miles (40 km) east of downtown Portland on the western slopes of the Cascade Range. The BRW is a surface water collection system, so the risk of landslide impact directly to the water and the infrastructure is relatively high. Because landslides are one of the most widespread and damaging natural hazards in Oregon, it is important to map and assess the risk in the BRW. The purpose of this study is to assist the Portland Water Bureau in understanding the landslide hazard better and thus increase their ability to reduce future risk. The study publication consists of a text report, five map plates, and three geodatabases.

EXPLANATION

This map depicts susceptibility to deep landslides in this area. For the purpose of this map, deep landslides are defined as those with a depth to the failure plane of greater than 15 ft (4.5 m) (Burns and Madin, 2009).

This susceptibility map was prepared by combining three factors: 1) landslide inventory data taken from the corresponding inventory map (see Plate 1) and head scarp buffers, 2) minimal zone buffer, and 3) geologic factors (susceptible geologic units and contacts, slope angles, and preferred direction of movement). The susceptibility zones resulting from the combination of these factors are shown in the Hazard Zone Matrix and shown in the Hazard Zone Matrix below. The deep landslide susceptibility data are displayed on top of a base map that consists of the lidar-derived digital elevation model. For additional detail on how this map was developed see the accompanying text report.

DEEP LANDSLIDE SUSCEPTIBILITY CLASSIFICATION

Each landslide susceptibility hazard zone shown on this map has been developed according to a classification scheme using a number of specific factors. The classification scheme was developed by the Oregon Department of Geology and Mineral Industries. The symbology used to display these hazard zones is explained below.

Deep Landslide Susceptibility Zone: This map uses color to show the relative degree of hazard. Each zone is a combination of several factors (see Hazard Zone Matrix, below).

HIGH: High susceptibility to deep landslides.
Deposits mapped as historical and/or active are outlined in black.

MODERATE: Moderate susceptibility to deep landslides.

LOW: Low susceptibility to deep landslides.

Deep Landslide Susceptibility Hazard Zone Matrix

Contributing Factors	Final Hazard Zone		
	High	Moderate	Low
1 Landslides, Head Scarp-Flanks, Buffers	included	—	—
2 Geologic Factors and Minimal Zone Buffer	—	included	—
3 Inferred Geologic Factors	—	—	included

1 Landslide Inventory

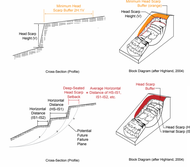
An inventory of all existing landslides in this area is shown on Plate 1. We prepared this inventory map by compiling all previously mapped landslides from published and unpublished geologic and landslide mapping, analyzing lidar-based geomorphology, and reviewing aerial photographs. We also attributed each landslide with classifications for activity, depth of failure, movement type, and confidence of interpretation. We created the inventory by using the protocol developed by Burns and Madin (2009). We extracted the deep landslides from the inventory and used them to create this deep landslide susceptibility map.

2 Head Scarp Buffers

Buffers were applied to all head scarps from the landslide inventory (Plate 1). In most cases the first buffer results in a minimum buffer distance and the second buffer (described below) results in the maximum buffer distance. In all cases the greater of the two was used.

The first buffer (orange on diagram) consists of a 2:1 horizontal to vertical distance (2H:1V). This buffer is different for each head scarp and is dependent on head scarp height. For example, a head scarp height of 18 ft (6 m) has a 2H:1V buffer equal to 36 ft (11.5 m).

The second buffer (red on diagram) is different for each head scarp and is dependent on the average of the horizontal distance between internal scarps. For example, an average horizontal distance of 150 ft (50 m) has a 2H:1V buffer equal to 300 ft (100 m).



3 Geologic Factors and Buffers

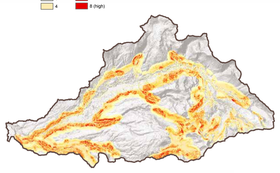
The small map below displays the scores of the relative geologic susceptibility zone factors, aminimal zone buffer applied around the high susceptibility zone, and the mapped deep landslide deposits in this area. The geologic zone factors are:

1) Susceptible geologic units
2) Susceptible geologic contacts
3) Susceptible slope angles for each engineering geology unit polygon
4) Susceptible direction of movement for each engineering geology unit polygon

The geologic susceptibility zone factors and the minimal zone buffer datasets along with professional judgment were used to create the boundary between the moderate and low deep landslide susceptibility zones.

A minimal zone buffer was applied around the high susceptibity zone of each landslide deposit. This buffer is different for each landslide deposit and is dependent on head scarp height.

Each geologic zone factor was given a score of 0, 1, or 2. Thus, if all factors have the highest score at some particular location, the final factor score is 8. A minimal combined factor score threshold between 3 and 5 along with professional judgment was used to delineate the boundary between the low and moderate zones.



LIMITATIONS

The deep landslide susceptibility map was developed following an established protocol (Burns and others, 2012) that incorporates several types of data. Several limitations are worth noting and underscore that any regional hazard map can be useful for regional applications but should not be used as an alternative to site-specific studies in critical areas. Limitations include the following.

1) Event-effect has been made to ensure the accuracy of the GIS and tabular database, but it is not feasible to completely verify all of the original input data.

2) As discussed in the above, the protocol to develop deep landslide susceptibility maps is based on three factors: 1) landslide inventory data taken from the corresponding inventory map and head scarp buffers, 2) minimal zone buffer, and 3) geologic factors (susceptible geologic units and contacts, slope angles, and preferred direction of movement). All of these parameters can affect the level of detail and accuracy of the final susceptibility map. Because the maps are based on a combination of factors, all of which have inherent uncertainty, the resultant hazard zones also have uncertainty.

a. Limitations of the landslide inventory, as discussed by Burns and Madin (2009).

b. Calculation of head scarp buffers is limited based on the head scarp height (first buffer) and an average of the horizontal widths of previous or downslope blocks (second buffer). It is assumed that all future deep landslides have the potential to fail retrogressively upslope; however, this is not always the case.

c. The additional factors used to delineate the moderate susceptibility zone include susceptible geologic units, susceptible geologic contacts, susceptible slope angles for each engineering geologic unit polygon, and susceptible direction of movement for each engineering geologic unit polygon. These factors are combined and a final score is produced, but the delineation of the final moderate zone is based on visual overlap of these four factors; therefore, the accuracy and resolution of the output data can be overestimated or underestimated.

3) The susceptibility maps are based on the topographic and landslide inventory data available as of the date of publication. Future landslides may render this map locally inaccurate.

4) The lidar-based digital elevation model does not distinguish elevation changes that may be related to anthropogenic features, such as roads, buildings, and dams. We used extensive GIS and field work to locate all of these existing structures and remove them or adjust the statistical properties in the model, such features have been included as a conservative approach and therefore must be examined on a site-specific basis.

5) Some landslides in the inventory may have been mitigated, thereby reducing their level of susceptibility. Because it is not feasible to collect detailed site-specific information on every landslide, potential mitigation has been ignored.

REFERENCES

Burns, W.J., and Madin, I.P., 2009, Protocol for inventory mapping of landslide deposits from light detection and ranging (lidar) imagery: Oregon Department of Geology and Mineral Industries Special Paper 42, 30 p., geodatabase template.

Burns, W.J., Mickelson, K.A., Jones, C.B., Pickner, S.G., Hughes, K.L., and Sloeter, R., 2013, Landslide hazard and risk study of northwestern Clackamas County, Oregon: Oregon Department of Geology and Mineral Industries, Open-File Report O-13-08, 38 p., 74 pl., scale 1:50,000 and 1:8,000, geodatabases.

Highland, L., compiler, 2004, Landslide types and processes: U.S. Geological Survey Fact Sheet 2004-3072 (ver. 1.1), 4 p.

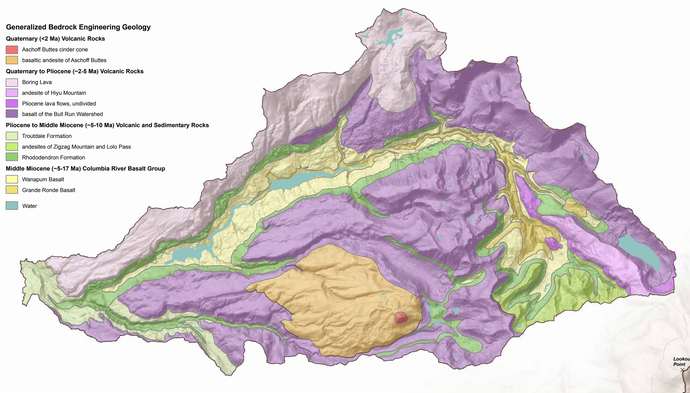
Generalized Bedrock Engineering Geology

Quaternary (<2 Ma) Volcanic Rocks
- Aschoff Buttes cinder cone
- basaltic andesite of Aschoff Buttes

Quaternary to Pliocene (~2-6 Ma) Volcanic Rocks
- Boring Lava
- andesite of Hiyu Mountain
- Pliocene lava flows, undivided
- basalt of the Bull Run Watershed

Pliocene to Middle Miocene (~5-10 Ma) Volcanic and Sedimentary Rocks
- Troutdale Formation
- andesites of Zigzag Mountain and Lolo Pass
- Rhododendron Formation

Middle Miocene (~5-17 Ma) Columbia River Basalt Group
- Wanapum Basalt
- Grande Ronde Basalt
- Water


BULL RUN WATERSHED BOUNDARY
Palmer Peak
Palmer Lake
Laboure Peak
Mount Talapus
Eagle Butte
Shorrard Point
Larch Mountain
Sandy River
Bull Run River
North Fork Bull Run River
Lookout Peak
Bear Creek House
Bull Run Dam Number 1
Bull Run Reservoir Number One
Big Bend Mountain
Table Mountain
Bull Run Lake
Buck Point
Walker Peak
Walker Prairie
Bull Run Reservoir Number Two
Headworks
Bull Run Dam Number 2
Preachers Peak
Sentinel Peak
Hiyu Mountain
Deer Meadows
Thunder Rock
Hoover Alder Butte
Township Meadow
Thimble Mountain
Halfway Hill
Burnt Peak
Aschoff Buttes
Yocum Lake
Hickman Lake
Hickman Butte
North Mountain
Clear Fork Bull Run
Little Sandy River
Derr Creek
Cedar Creek
Deer Creek
Bear Creek
Boulder Creek
Dairy Creek
Devil Creek
Trout Creek
Cedar Spring Creek
Bonney Creek
Fairy Creek
Lost Creek
Little Sandy Creek
Roaring Creek
Alder Creek
Hunchback Creek
Fifteenmile Creek
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